

JURNAL METAMORFOSA Journal of Biological Sciences elSSN: 2655-8122 http://ojs.unud.ac.id/index.php/metamorfosa

Umur, Lingkungan Pengendapan, dan Karakteristik Fosil Palinomorf Formasi Batuasih, Sukabumi, Jawa Barat

Age, Depositional Environment, and Characteristics of Palynomorph Fossil in the Batuasih Formation, Sukabumi, West Java

Rizki Satria Rachman¹*, Winantris², Budi Muljana³

^{1,2,3)} Geological Engineering, Faculty of Geological Engineering, Padjadjaran University, Jl.Raya Bandung – Sumedang Km.21, Sumedang Regency, West Java, 45363, Indonesia *Email: rizkisatriarachman@gmail.com

INTISARI

Palinomorf merupakan sisa material organik dari makhluk hidup yang berukuran debu dengan sifat tahan terhadap zat asam kuat yang dapat mencerminkan bagaimana kondisi umur dan lingkungan saat suatu batuan terbentuk. Formasi Batuasih merupakan salah satu formasi batuan di Cekungan Bogor berumur tua yang masih sangat sedikit diinterpretasikan terutama dari aspek palinomorfnya. Penelitian ini bertujuan untuk melihat karakteristik palinomorf, umur relatif, dan lingkungan dari batuan di Formasi Batuasih. Metode penelitian meliputi kegiatan lapangan dengan melakukan pengukuran penampang stratigrafi terukur dan sampling batuan dengan interval 5 meter atau setiap perubahan litologi, kegiatan laboratorium meliputi kegiatan preparasi sampel dengan metode asam treatment, identifikasi dilakukan menggunakan mikroskop binocular CX-22, dan analisis umur dan lingkungan pengendapan ditarik berdasarkan asosiasi palinomorf yang hadir pada batuan. Hasilnya, Palinomorf hadir pada 8 dari 10 sampel diantaranya (B.1, B.2, B.4, B.6 – B.10). 169 individu palinomorf teridentifikasi terdiri dari 19 spesies yang berbeda. Proxapertites merupakan genus yang mendominasi baik itu Proxapertites operculatus, Proxapertites cursus, dan Proxapertites psilatus. Sedangkan spora didominasi datang species Verrucatosporites usmensis. Selain polen dan spora, ditemukan juga palinomorf berupa dinoflagellata dan foraminifera test lining. Analisis palinomorf mencerminkan Formasi Batuasih memiliki umur Eosen akhir – Oligosen awal (41, 2 - 27, 8 jtl) dengan lingkungan pengendapan transisi.

Kata kunci: Palinomorf, Formasi Batuasih, Karakteristik, Umur, Lingkungan

ABSTRACT

Palinomorphs are the remains of organic material from living things that are dust-sized with properties that are resistant to strong acids, which can reflect how the age and environmental conditions were when a rock was formed. Batuasih Formation is one of the oldest rock formations in Bogor Basin which is rarely interpreted, especially from its palynological aspect. This study aims to see palynomorph characteristics, relative age, and environmental conditions in Batuasih Formation. There are several steps in conducting the research. Field activities by measured stratigraphic sections and sampling every 5-meter intervals or each lithological changes, laboratory activities by sample preparation using acid treatment method, identification were carried out using CX-22 binocular microscope, and laboratory analysis of age and depositional environment is drawn based on palynomorph associations that are present in rocks. The result found that palynomorphs were present in 8 of 10 samples (B.1, B.2, B.4, B.6 - B.10). 169

palynomorphs identified to consist of 19 different species. Dominant pollen came from genus *Proxapertites*, both *Proxapertites operculatus*, *Proxapertites cursus*, and *Proxapertites psilatus*. While dominant spores came from *Verrucatosporites usmensis*. Apart from pollen and spores, Palynomorphs in the form of dinoflagellates and foraminifera test lining are also present in this formation. Palynomorph analysis shows that Batuasih Formation has Late Eocene - Early Oligocene age (41.2 - 27.8 Ma) with transitional depositional environment.

Keywords: Palynomorph, Batuasih Formation, Characteristics, Age, Environment

INTRODUCTION

Java Island is part of Indonesia which is located in the primary Indo-Pacific environment and has very high rainfall and tropical climate (Dubois *et al.*, 2014). Its existence in the tropical region makes Java Island has a very high diversity of flora called mega-biodiversity. This diversity has continued to change, especially since the last ice age occurred in Indonesia (Kaars & Dam, 1995; Kaars & Dam, 1997; Kaars *et al.*, 2001; Wibowo *et al.*, 2016). Flora always has a desire to reproduce, one method of plant reproduction is through pollination of pollen and spores (Ritchie & Lichti-Federovich, 1967; Di-Giovanni & Kevan, 1991).

Palynomorph is a dust-sized material that resistant to strong acids (Riding & Kyffin-Hughes, 2004; Traverse, 2007; Dutta et al., 2013). Pollen and spores are the main components of palynomorphs in addition to other materials such as algae, acritarch, and foraminifera lining test (Miller et al., 1982; Baioumi et al., 2012). During the reproduction process, pollen and spores can fall to the surface and preserved to become fossils in the rock (Slater and Wellman, 2015; Xu et al., 2016). The presence of palynomorphs in these rocks can reflect the age and environmental conditions when these rocks formed (Hope et al., 2004; Winantris et al., 2006; Setijadi et al., 2015). Seeing palynomorphs have wide distribution and complete evolution, they are suitable for determining relative age and environmental

conditions of rocks, especially in terrestrial and transitional environments (Hillen, 1986; Larsson *et al.*, 2006; Macphail & Cantrill, 2006).

Research using palynomorphs to determine age and environment sediment has been widely carried out in Java region, but most of these studies have focused on Holocene age (Kaars & Dam, 1995; Kaars & Dam, 1997; Kaars et al., 2001). Batuasih Formation is one of the oldest formations in Bogor Basin (Figure 1d) with Early Oligocene - Late Oligocene age interval (33.9 - 23.03 Ma) which interpreted to be deposited in the transitional environment shallow seas from previous studies (Baumann, 1972; Effendi et al., 1998; Martodjodjo, 2003; Clements, 2007; Hendrizan, 2012; Wibowo & Kapid, 2014). Research on fossils in Batuasih Formation has been very few, especially when viewed from the palynological aspect. The research location is in Batuasih Formation which exposed in Cibatu River area, Sukadamai Village, Sukabumi Regency, West Java with coordinates 06°56'29.64" South Latitude -106°50'30.18" East Longitude to 06°56'37.47" South Latitude - 106°50'32.95" East Longitude (Figure 1). From this explanation, it becomes interesting to see how palynomorph content, relative age, and environmental conditions of Batuasih Formation rocks if interpreted through its Palynological aspect.



Figure 1. Research area; a. Research location is related to West Java; b. Research location is related to the Walat, Batuasih, and Rajamandala Formations (Effendi *et al.*, 1998); c. Sampling and measured stratigraphic section results; d. Stratigraphy of study area

MATERIALS AND METHODS

The research method is divided into several steps. These steps include field activity, laboratory activity in the form of sample preparation, and laboratory analysis in the form of palynomorph characteristics identification, age, and depositional environment analysis.

Field activity was carried out by measured stratigraphic sections of Batuasih Formation rocks which exposed in Cibatu River (Siregar, 2005; Fauzi, 2017; Subagja *et al.*, 2019). Sampling mainly focuses on any lithological changes, but if thickness of the rock layer exceeds 5 meters, rock sampling takes at 5-meter intervals. The sample then prepared in the laboratory. Laboratory preparation was conducted by acid treatment to separate palynomorph material from other materials using

certain chemicals (Grey, 1999; Azizah *et al.*, 2016; Fajrina *et al.*, 2016; Yulianto *et al.*, 2019). Preparation results then identified using CX-22 binocular microscope to see the characteristics of pollen and spores present in Batuasih formation (Winantris *et al.*, 2012; Marcos *et al.*, 2015).

In this study, the age and depositional environment of Batuasih Formation were identified by looking at palynomorph content in these rocks. Age and depositional environment are drawn based on palynomorph associations present in rocks which represent the relative age and environmental conditions at the time Batuasih Formation was formed (McGlone *et al.*, 1978; KSSI, 1996; Polhaupessy, 2009; Yulianto *et al.*, 2019).

Fossil Name	Sample											F unction and the second seco
	B.1	B.2	B.3	B.4	B.5	B.6	B.7	B.8	B.9	B.10	Т	Environment
Acrostichum sp.				2		1	1				4	Fluvial-lacustrine, Back mangrove
Chepalomappa malloticarpa				3		2	2				7	Freshwater swamp
Clavifera triplex						1					1	Transitional
Crassoretitriletes vanraadshoveni	1					1					2	Freshwater swamp
Dicolpopollis sp.		1		1		1			1		4	Freshwater swamp
Distaeverrusporites simplex									1		1	Tidak Signifikan
Foraminifera Test Lining							2				2	Marine
Laevigatosporites						2	1		4		7	Fluvial-deltaic, Lacustrine, Freshwater
Lycopodium cernuum							1				1	Fluvial-lacustrine
Palmaepollenites kutchensis						1	1	1			3	Mangrove, Back-mangrove
Podocarpidites marwickii		1								1	2	Montane-forest
Proxapertites cursus						1	2		2		5	Fluvial-deltaic
Proxapertites operculatus		1		18		24	27	9	22	1	102	Fluvial-deltaic
Proxapertites psilatus				1		6	4		2		13	Fluvial-deltaic
Spiniferites pseudofurcatus							1				1	Marine-Lagoon
Spiniferites ramosus				1							1	Marine-deltaic
Trichotomosulcites subgranulatus							1				1	Fluvial
Tricolpites confessus				1			1				2	Fluvial
Verrucatosporites usmensis				1		3	5		1		10	Freshwater swamp
Total	1	3	0	28	0	43	49	10	33	2	169	

Table 1. Palinomorph Content of Batuasih Formation

RESULTS

Research area has a lithology which is dominated by mudstone which in some places alternates with sandstone and limestone. The thickness of rock outcrop that was found reached 74 meters on the edge of Cibatu River. The characteristics of mudstone are massive, dark gray color with pyrite content in this rock. Sandstone has a massive light gray color with fine sand grain size. Meanwhile, limestone has white and hard characteristics. From this rock outcrop samples were taken for palynomorph analysis (Figure 1c).

Palynomorphs found in Batuasih Formation are very diverse. From 10 samples (B.1 - B.10), palynomorphs were present in 8 samples (B.1, B.2, B.4, B.6 - B.10). The number of palynomorphs observed reached up to 169 individuals which identified into 19 different species. The most dominant pollen palynomorphs come from the genus Proxapertites, both Proxapertites operculatus, *Proxapertites cursus*, and *Proxapertites psilatus*.

While spores in Batuasih Formation are dominated by *Verrucatosporites usmensis*. Apart from pollen and spores, palynomorphs in Batuasih Formation sample also found other palynomorphs, and this palynomorph is in the form of dinoflagellates and foraminifera test lining (Table 1).

Palynomorph Characteristics

The characteristics of pollen and spores (palynomorphs) in Batuasih Formation are very numerous in terms of size, shape, aperture/laesure, and ornamentation. The following describes each palynomorph found in the Batuasih Formation.

Fossil Name : Acrostichum sp.

Botanical Affinity : Acrostichum

Description (Figure 2e) :

Spores with size $30 - 50 \mu$ (Medium); Figure 50 x 47 μ ; Unity monad, Brownish yellow color, Polarity polar, Distal axis rounded; Proximal axis slightly pointed; Shape Amb triagular;

Ekuatorial subspheroidal (P/E = 6/8 - 8/6), Laesure trilete L = 12μ ; W = thin < 1μ ; Laesura extends 2/3 center to surface, Ornamentation psilate – locally scabrate < 1μ , Intectate, and Heteropolar.

Fossil Name : *Chepalomappa malloticarpa* Botanical Affinity : *Chepalomappa malloticarpa* Description :

Pollen with size $30 - 60 \mu$ (Medium - Large); Figure 61 x 43 μ , Unity monad, Brownish yellow color, Polarity oblique, Shape Amb circular / rounded; Ekuatorial peroblate - oblate (P/E = < 4/8 - 6/8), Aperture inaperture difficult to observe, Ornamentation reticulate; Muri > 1 μ ; Lumen > 2 μ ; Brochi not uniform circular polygonal, Intectate, and Isopolar.

Fossil Name: Clavifera triplex

Botanical Affinity : Gleicheniaceae Description :

Spores with size $25 - 40 \mu$ (Medium); Figure 32 x 25 μ , Unity monad; Brownish yellow color, Polarity polar, Shape Amb triangular concave; Ekuatorial subspheroidal (P/E = 6/8 - 8/6), Laesure trilete L = 5 μ ; W = thin < 1 μ ; Wavy extending from the center to the surface; There was thickening of columella on surface and center, Ornamentation psilate, Tectate, and Heteropolar.

Fossil Name : *Crassoretitriletes vanraadshooveni* Botanical Affinity : *Lygodium microphyllum*

Description (Figure 2f) :

Spores with size $58 - 101 \mu$ (Large – Very large); Figure $58 \times 54 \mu$, Unity monad, Brown color, Polarity oblique, Distal axis rounded; Proximal axis slightly pointed; Shape Amb triagular convex – sub-circular; Ekuatorial subspheroidal (P/E = 6/8 – 8/6), Laesure trilete L = 15 μ ; W = thin; laesure covered in ornamentation, Ornamentation reticulate; Muri 2 μ ; Lumen > 2 μ ; Brochi uniform circular - polygonal, Tectate, and Heteropolar.

Fossil Name : *Dicolpopollis* sp. Botanical Affinity : Palmae Description (Figure 2a) : Pollen with size $24 - 50 \mu$ (Medium); Figure 50 x 27 μ , Unity monad, Brownish yellow color, Polarity polar, Shape Amb elliptical; Ekuatorial oblate (P/E = 4/8 - 6/8), Aperture disulcate (2-Sulcate) L = 15μ ; W = 3μ ; thinning towards center; Longicolpate until merge, Ornamentation reticulate; Muri 1 μ ; Lumen 1 μ ; Brochi uniform rectangular - polygonal, Tectate, and Isopolar.

Fossil Name : *Distaeverrusporites simplex* Botanical Affinity : Unknown Description :

Spores with size $30 - 40 \mu$ (Medium); Figure 37 x 37 μ , Unity monad, Orange color, Polarity oblique, Shape Amb circular / rounded; Ekuatorial subspheroidal (P/E = 6/8 - 8/6), Laesure alete difficult to observe covered in ornamentation, Ornamentation reticulate; Muri > 1 μ ; Lumen > 1 μ ; Brochi uniform circular - polygonal, Intectate, and Isopolar.

Fossil Name : *Foraminifera Test Lining* Botanical Affinity : *Foraminifera* Description (Figure 2j) :

Foraminifera with size $50 - 70 \mu$ (Large); Figure 69 x 58 μ , Unity monad, Brown color, Polarity oblique; Shape Amb sub-circular; Ekuatorial subspheroidal (P/E = 6/8 - 8/6), Divided into chamber; Planispiral, Chamber L = 21 μ and W = 20 μ ; Suture 4 μ , Ornamentation psilate,

Fossil Name : Laevigatosporites

Intectate, and Isopolar.

Botanical Affinity : Polypodiaceae Description (Figure 2g) :

Spores with size $30 - 60 \mu$ (Medium - Large); Figure $60 \times 38 \mu$, Unity monad, Brownish yellow color, Polarity ekuatorial, Shape Amb peanut / sub-circular; Ekuatorial peroblate - oblate (P/E = < 4/8 - 6/8), Laesure monolete L = 31μ ; W = 3μ ; Laesure is thickened in the middle and extends but not to the surface of spores, Ornamentation psilate, Intectate, and Heteropolar.

Fossil Name : *Lycopodium cernuum* Botanical Affinity : *Lycopodium cernuum* Description :

Spores with size $30 - 60 \mu$ (Medium - Large); Figure $38 \times 32 \mu$, Unity monad, Brownish yellow color, Polarity polar, Shape Amb triagular; Ekuatorial subspheroidal (P/E = 6/8 - 8/6), Laesure trilete L = 15μ ; W = thin thickened to the center until 2μ ; Laesure extends 1/2 center to surface, Ornamentation fine-rugulate; Thickening at edge of the spores, Intectate, and Heteropolar.

Fossil Name : *Palmaepollenites kutchensis* Botanical Affinity : *Palmae* Description :

Spores with size $27 - 152 \mu$ (Medium – Very large); Figure $38 \times 37 \mu$, Unity monad, Brownish yellow color, Polarity polar, Shape Amb peanut / sub-circular; Ekuatorial oblate - subspheroidal (P/E = 4/8 - 8/6), Laesure monolete L = 35μ (Half Individual); W = 11μ ; Laesura is thickened in middle and extends to the surface of the spores, Ornamentation psilate, Intectate, and Heteropolar.

Fossil Name : *Podocarpidites marwickii* Botanical Affinity : *Podocarpus*

Description (Figure 2b) :

Pollen with size corpus $37 - 51 \mu$ (Medium -Large) and saccate $17 - 30 \mu$ (Medium); Figure $49 \times 40 \mu$, Unity monad, Yellow color, Polarity polar, Shape Amb circular / bisaccate; Ekuatorial peroblate - oblate (P/E = < 4/8 - 6/8), Inaperture bisaccate (2-Saccate); Corpus L = 37μ ; W = 26μ ; Saccate L = 33μ ; W = 24μ ; Surface to saccate = 3μ ; Between saccate = 1μ , Ornamentation reticulate - psilate; Saccate reticulate; Muri < 1μ ; Lumen 1 μ ; Brochi uniform polygonal; Corpus granulate or scabrate; Ornamentation uniform, Tectate, and Heteropolar.

Fossil Name : Proxapertites cursus

Botanical Affinity : Araceae or Arecaceae Description :

Pollen with size $40 - 60 \mu$ (Medium - Large); Figure 55 x 35 μ , Unity monad, Brownish yellow color, Polarity oblique, Flat on equatorial view; Shape Amb circular / rounded; Ekuatorial subspheroidal (P/E = 6/8 - 8/6), Aperture monosulcate difficult to observe; Irregular; Aperture extends along surface of the pollen in equatorial part, Ornamentation reticulate; Muri > 1 μ ; Lumen 1 μ ; Brochi uniform polygonal increase irregular approaching the sulcus, Tectate, and Heteropolar.

Fossil Name : *Proxapertites operculatus* Botanical Affinity : Araceae or Arecaceae Description (Figure 2c):

Pollen with size $40 - 60 \mu$ (Medium - Large); Figure 52 x 37 μ , Unity monad, Brownish yellow color, Polarity oblique, Flat on equatorial view; Shape Amb circular / rounded; Ekuatorial subspheroidal (P/E = 6/8 - 8/6), Aperture monosulcate L = 42 μ ; W = 7 μ ; Irregular; Aperture extends along surface of the pollen in equatorial part, Ornamentation fine-reticulate; Muri < 1 μ , Lumen < 1 μ ; Brochi uniform circular - polygonal, Tectate, and Heteropolar. Fossil Name : *Proxapertites psilatus*

Botanical Affinity : Araceae or Arecaceae Description (Figure 2d):

Pollen with size $40 - 60 \mu$ (Medium - Large); Figure $60 \times 38 \mu$, Unity monad, Brownish yellow color, Polarity oblique, Flat on equatorial view; Shape Amb circular / rounded; Ekuatorial subspheroidal (P/E = 6/8 - 8/6), Aperture monosulcate difficult to observe; Irregular; Aperture extends along surface of the pollen in equatorial part, Ornamentation psilate, Intectate, and Heteropolar.

Fossil Name : *Spiniferites pseudofurcatus* Botanical Affinity : *Spiniferites pseudofurcatus* Description (Figure 2i):

Dinoflagellates with size $30 - 60 \mu$ (Medium -Large); Figure 43 x 36 μ , Unity monad, Yellow color, Polarity oblique, Shape Amb circular / rounded; Ekuatorial subspheroidal (P/E = 6/8 -8/6), Ornamentation in the form of arms L = 5-9 μ and W = thin < 1 μ ; The arms is split into 2 parts to form "v" at the ends, Ornamentation psilate, Intectate, and Isopolar.

Fossil Name : *Spiniferites ramosus* Botanical Affinity : *Spiniferites ramosus* Description :

Dinoflagellates with size $70 - 90 \mu$ (Medium -Large); Figure 88 x 75 μ , Unity monad, Brownish yellow color, Polarity oblique, Shape Amb circular / rounded; Ekuatorial subspheroidal (P/E = 6/8 - 8/6), Ornamentation in the form of arms $L = 16 \mu$ and $W = 1-2 \mu$; The arms is split in 2 parts at the end to form an

elbow, Ornamentation psilate, Intectate, and Isopolar.



Figure 2. Palynomorph Batuasih Formation; a-d. Pollen; a. Dicolpopollis sp.; b. Podocarpidites marwickii; c. Proxapertites operculatus; d. Proxapertites psilatus; e-h. Spores; e. Acrostichum sp.; f. Crassoretitriletes vanraadshoveni; g. Laevigatosporites; h. Verrucatosporites usmensis; i. Dinoflagellates; i. Spiniferites pseudofurcatus; j. Foraminifera Test Lining; Scale Bar = 10 μ

Fossil Name : *Trichotomosulcites subgranulatus* Botanical Affinity : Dicotyledonae Description :

Pollen with size $31 - 47 \mu$ (Medium); Figure 36 x 32 μ , Unity monad, Brownish yellow color, Polarity oblique, Shape Amb triangular - subcircular; Ekuatorial subspheroidal (P/E = 6/8 - 8/6), Inaperture Saccate L (Surface to triangle) = 11 μ ; W (Triangle) = 10 μ ; Saccate forms triangle at the center of the pollen, Ornamentation psilate - microgranulate - microreticulate; Muri 1 μ ; Lumen 1 μ ; Brochi uniform circular - polygonal, Intectate, and Isopolar.

Fossil Name : *Tricolpites confessus* Botanical Affinity : *Tricolpites confessus* Description :

Pollen with size $25 - 50 \mu$ (Medium); Figure 47 x 44 μ , Unity monad, Brownish yellow color, Polarity polar, Shape Amb circular / rounded / lobate convex; Ekuatorial subspheroidal (P/E = 6/8 - 8/6), Aperture tricolpate (3-Colpate) L = 9 μ ; W = 14 μ ; Longicolpate extends from the surface to center < $\frac{1}{2}$ body, Ornamentation fine-reticulate; Muri < 1 μ ; Lumen < 1 μ ; Brochi uniform polygonal, Punctectate, and Isopolar.

Fossil Name : *Verrucatosporites usmensis* Botanical Affinity : *Stenochlaena palustris* Description (Figure 2h) :

Spores with size $39 - 61 \mu$ (Medium - Large); Figure 57 x 33 μ , Unity monad, Brown color,

Polarity ekuatorial, Bilateral symmetry; Distal axis convex; Proximal axis straight until slightly concave; Shape Amb peanut / sub-circular; Ekuatorial peroblate - oblate (P/E = < 4/8 - 6/8), Laesure monolete covered in ornamentation, Oenamentation gemmate L & W = 3-10 µ; Size and shape of ornamentation is not uniform and decreasing towards the laesura, Tectate, and Heteropolar.

Age and Depositional Environment Batuasih Formation

In this study, Batuasih Formation has Late Eocene to Early Oligocene age range (41.2 - 27.8 Ma). The age range is drawn based on biostratigraphic interval zone between the presence of *Proxapertites* group which became extinct in Late Eocene and *Crassoretitriletes vanraadshooveni* that only emerged at Early Oligocene age (Figure 3a). Additionally, the presence of genus *Podocarpus* (*Podocarpidites*) indicates that the relative age of Batuasih Formation is in Eocene - Oligocene boundary (Morley, 2010; Morley, 1991).

Depositional environment of Batuasih Formation in this study is in a transitional environment, defined by the presence of large numbers of transitional palynomorphs such as Acrostichum and Palmaepollenites kutchensis. The dominance of Proxapertites group can also be a key indicator of transitional depositional environment. This Proxapertites group can be present along rivers (Fluvial) to transitional parts, either deltas or lagoons (Germeraad et al., samples found 1968). Some marine palynomorphs such as dinoflagellates and foraminifera test lining. It indicates that Batuasih Formation is no longer in a terrestrial environment but is more vital towards the transitional environment (Vernal *et al.*, 1993; Sawada *et al.*, 1999; Raine & Askin, 2001).

The depositional environment interpreted by palynomorphs cannot be directly drawn based on the dominance of palynomorphs. Instead, an interpretation must be drawn based on the presence and association of plants reflected by the presence of palynomorphs in a rock sample. Besides, palynomorphs from terrestrial can become numerous in the transition region because these palynomorphs are transported and deposited different environments in (downstream direction). Therefore. the depositional environment interpretation was carried out by looking at the presence of palynomorphs and their associations (Haseldonckx, 1974; Haseldonckx, 1977).

It can be seen from the diagram that as moving towards the younger sample, the terrestrial palynomorphs are getting more replaced by the transitional palynomorphs. It indicates a more substantial transition effect towards Batuasih Formation sample (Figures 3b - 3c). In some samples, it appears that palynomorph terrestrial (Fluvial) are dominant due to differences in the number of individuals of each sample. However, if seen from the total percentage of each environment, it can be seen that the effect of the transition palynomorph is very high, reaching up into 43% (Figure 3d).



Figure 3. Age and deposition environment of the Batuasih Formation; a. Age range from interval zone of Batuasih Formation; b. Palynomorph distribution in each depositional environment; c. Comparison of fluvial and transition palynomorphs of study samples; d. Legend and percentage of palynomorphs for each environment

DISCUSSION

Previous research states that Batuasih Formation was formed at Early Oligocene - Late Oligocene age (33.9 - 23.03 Ma) which interpreted through other objects such as foraminifera and nannoplankton. (Baumann, 1972; Effendi et al., 1998; Martodjodjo, 2003; Clements, 2007; Hendrizan, 2012; Wibowo & Kapid, 2014). From this, it can be seen that the age of Batuasih Formation has an older age (Late Eocene). This Late Eocene age cannot be neglected because key palynomorphs of Late Eocene age in Batuasih Formation are the dominant palynomorphs. This dominant palynomorph is **Proxapertites** group (*Proxapertites* cursus, **Proxapertites** operculatus, and Proxapertites psilatus). While the last age of Batuasih Formation found in this study was limited to Early Oligocene. It happened because Proxapertites group became extinct at this age (Morley, 1991). Moreover, the

presence of *Podocarpidites* indicates that Batuasih Formation has relative age not far from Eocene - Oligocene boundary (33.9 Ma). It happened because Podocarpidites appeared in Indonesia in the Late Eocene. Podocarpidites came to Indonesia after spreading from India and Ninety East Ridge at the Paleocene age (Morley, 2010; Morley, 1991). Another source states that coniferous groups in Indonesia were present after Late Eocene age. It occurred in association with Late Eocene collision between the Indian plate and the Eurasian plate (Dupont-Nivet et al., 2013). Therefore, the presence of *Podocarpidites* indicates that the age of Batuasih Formation is not far from the Eocene - Oligocene boundary. Other studies that extend the age of Batuasih Formation to Late Oligocene can occur. It happened because Batuasih Formation was exposed in other areas. Therefore, there is the possibility of finding Batuasih Formation, which has Late Oligocene age in that region.

From this interpretation, the depositional environment of this research is consistent with results of previous studies. Previous research has interpreted Batuasih Formation to form in the transitional environment to shallow seas 1972; Effendi et (Baumann, al., 1998; Martodjodjo, 2003; Clements, 2007; Hendrizan, 2012; Wibowo & Kapid, 2014). As for this study, Batuasih Formation is interpreted to form in a transitional environment characterized by the presence of marine palynomorphs and the strong influence of transition palynomorphs in the study sample.

CONCLUSION

Batuasih Formation is one of the rock formations in Bogor Basin that shows the presence of palynomorphs. As a result, 169 individuals palynomorph from 19 species were identified. Dominant pollen came from genus Proxapertites, both Proxapertites operculatus, Proxapertites cursus, and Proxapertites psilatus. While dominant spores come from Verrucatosporites usmensis. Palynomorphs in form of dinoflagellates and foraminifera test lining are also present in this formation. From palynomorph analysis, Batuasih Formation has Late Eocene - Early Oligocene age (41.2 - 27.8 ma) with the transitional depositional environment.

ACKNOWLEDMENT

Thank you to the Chancellor of Padjadjaran University for funding research through the HIU-RKDU program, Paleontology laboratory team of the Faculty of Geological Engineering UNPAD, and all parties who have helped and provided encouragement in this research.

REFERENCES

Azizah, N., Suedy, S. W. A., and Prihastanti, E.
2016. Keanekaragaman Tumbuhan Berdasarkan Morfologi Polen dan Spora dari Sedimen Telaga Warna Dieng, Kabupaten Wonosobo, Jawa Tengah, *Buletin Anatomi dan Fisiologi*, 24(1): 66-75.

- Baioumi, A. E. H. A., Mandur, M. M., and Moustfa, T. 2012. Aptian – Cenomanian Palynozonation and Paleoecology from Horous-1 Well, Northern Western Desert, Egypt, Journal of Applied Sciences Research, 8(3): 1490-1501.
- Baumann, P. 1972. The Cenozoic of Java and Sumatera, 1st Annual Convention Proceedings, Indonesian Petroleum Association, 1972, hal. 31-40.
- Clements, B. 2007. Cretaceous to Late Miocene Stratigraphic and Tectonic Evolution Of West Java, 31st Annual Convention Proceedings, Indonesian Petroleum Association, 2007, hal. 1-18.
- Di-Giovanni, F., and Kevan, P. G., 1991. Factors Affesting Pollen Dynamics and Its Importance to Pollen Contamination: A Review, *Canadian Journal of Forest Research*, 21(8): 1155-1170.
- Dubois, N., Oppo, D. W., Galy, V. V., Mohtadi,
 M., Kaars, S. V. D., Tierney, J. E.,
 Rosenthal, Y., Eglinton, T. I., Luckge, A.,
 and Linsley, B. K., 2014. Indonesian
 Vegetation Response to Changes in
 Rainfall Seasonality Over The Past 25,000
 Years, *Nature Geoscience*, 7: 513–517.
- Dupont-Nivet, G., Hoorn, C., and Konert, M. 2008. Tibetan Uplift Prior to The Eocene-Oligocene Climate Transition: Evidence from Pollen Analysis of The Xining Basin, *Geology*, 36: 987-990.
- Dutta, S., Hartkopf-Froder, C., Witte, K., Brocke, R., and Mann, U., 2013. Molecular Characterization of Fossil Palynomorphs by Transmission Micro-FTIR Spectroscopy: Implications for Hydrocarbon Source Evaluation, *International Journal of Coal Geology*, 115: 13-23.
- Effendi, Kusnama, and Hermantos. 1998. Peta Geologi Lembar Bogor, Jawa Barat, Bandung: Pusat Penelitian dan Pengembangan Geologi.
- Fajrina, D., Winantris, and Fauzielly, L. 2016.
 Sejarah Perubahan Iklim Berdasarkan
 Analisis Palinologi Daerah Derwati,
 Bandung, Jawa Barat, Prosiding Seminar
 Nasional MIPA, Universitas Padjadjaran,

Jatinangor 27-28 Oktober 2016, hal. 198-204.

- Fauzi, R. 2017. Studi Lingkungan Pengendapan Formasi Balikpapan Daerah Batu Besaung, Prosiding Seminar Nasional Teknologi, Inovasi dan Aplikasi di Lingkungan Tropis IV, Universitas Mulawarman, Samarinda 9 November 2017, hal. 76-80.
- Germeraad, J. H., Hoping, C. A., and Muller, J. 1968. Palynology of Tertiary Sediments From Tropical Areas, *Review of Palaeobotany and Palynology*, 6(3-4): 189-348.
- Grey, K. 1999. A Modified Palynological Preparation Technique for The Extraction of Large Neoproterozoic Acanthomorph Acritarchs and Other Acid Insoluble Microfossils, Perth: Western Australia Geological Survey.
- Haseldonckx, P. 1974. A palynological interpretation of palaeoenvironments in SE Asia, *Sains Malaysiana*, 3: 119-127.
- Haseldonckx, P. 1977. The Palynology Of A Holocene Marginal Peat Swamp, *Review* of Palaeobotany and Palynology, 24: 227-238.
- Hendrizan, M., Praptisih, and Putra, P. S. 2012.
 Depositional Environment of The Batuasih Formation on The Basis of Foraminifera Content: A Case Study in Sukabumi Region, West Java Province, Indonesia, *Indonesian Journal of Geology*, 7(2): 101-112.
- Hillen, R. 1986. Palynology as A Tool in Delineating Tropical Lowland Depositional Environments of Late Quaternary Age, *Geosea*, 1: 495-504.
- Hope, G., Kershaw, A. P., Kaars, S. V. D., Xiangjun, S., Liew, P., Heusser, L. E., Takahara, H., McGlone, M., Miyoshi, N., and Moss, P. T., 2004. History of Vegetation and Habitat Change in The Austral-Asian region, *Quaternary International*, 118-119: 103-126.
- Kaars, S. V. D., and Dam, R. 1997. Vegetation and Climate Change in West-java, Indonesia During The Last 135,000 Years, *Quaternary International*, 37: 67-71.

- Kaars, S. V. D., Penny, D., Tibby, J., Fluin, J., Dam, R. A. C., and Suparan, P., 2001. Late Quartenary Palaeoecology, Palynology, and Palaeolimnology of A Tropical Lowland Swamp: Rawa Danau, West Java, Indonesia, *Palaeogeography, Palaeoclimatology, Palaeoecology,* 171(3-4): 185-212.
- Kaars, W. A. V. D., and Dam, M. A. C., 1995. A 135,000-year Record of Vegetational and Climatic Change from The Bandung Area, West-Java, Indonesia, *Palaeogeography*, *Palaeoclimatology*, *Palaeoecology*, 117(1-2): 55-72.
- KSSI (Komisi Sandi Stratigrafi Indonesia). 1996. Sandi Stratigrafi Indonesia, Jakarta: Ikatan Ahli Geologi Indonesia.
- Larsson, L. M., Vajda, V., and Rasmussen, E. S. 2006. Early Miocene Pollen and Spores from Western Jylland, Denmark -Environmental and Climatic Implication, *GFF*, 128(3): 261-272.
- Macphail, M., and Cantrill, D. J. 2006. Age and Implications of The Forest Bed, Falkland Islands, Southwest Atlantic Ocean: Evidence from Fossil Pollen and Spores, *Palaeogeography, Palaeoclimatology, Palaeoecology*, 240(3-4): 602-629.
- Marcos, J. V., Nava, R., Cristobal, G., Redondo, R., Escalante-Ramirez, B., Bueno, G., Deniz, O., Gonzalez-Porto, A., Pardo, C., Chung, F., and Rodriguez, T., 2015. Automated Pollen Identification Using Microscopic Imaging and Texture Analysis, *Micron*, 68: 36-46.
- Martodjojo, S. 2003. Evolusi Cekungan Bogor, Jawa Barat, Bandung: Institut Teknologi Bandung.
- Miller, A. A. L., Mudie, P. J., and Scott, D. B. 1982. Holocene History Of Bedford Basin, Nova Scotia: Foraminifera, Dinoflagellate, And Pollen Records, *Canadian Journal of Earth Sciences*, 19(12): 2342-2367.
- Morley, R. J. 2010. Dispersal and Paleoecology of Tropical Podocarps, *Smithsonian Contributions To Botany*, 21-41.
- Morley, R. J. 1991. Tertiary stratigraphic palynology in Southeast Asia: current

status and new directions. *Geological Society of Malaysia*, 28: 1-36.

- Polhaupessy, A. A. 2009. Polen Paleogen-Neogen dari Daerah Nanggulan dan Karangsambung Jawa Tengah, *Jurnal Sumber Daya Geologi*, 19(5): 325-332.
- Raine, J. I., and Askin, R. A. 2001. Terrestrial palynology of Cape Roberts Project drillhole CRP-3, Victoria Land Basin, Antarctica, *Terra Antartica*, 8(4): 389-400.
- Riding, J. B., and Kyffin-Hughes, J. E. 2004. A Review of The Laboratory Preparation of Palynomorphs With A Description of An Effective Non-Acid Technique, *Revista Brasileira de Paleontologia*, 7(1): 13-44.
- Ritchie, J. C., and Lichti-Federovich, S. 1967. Pollen Dispersal Phenomena In Arctic-Subarctic Canada, *Review of Palaeobotany and Palynology*, 3: 255-266.
- Sawada, M., Gajewski, K., Vernal, A. D., and Richard, P. 1999. Comparison of marine and terrestrial Holocene climatic reconstructions from northeastern North America, *The Holocene*, 9(3): 267-277.
- Setijadi, R., Widagdo, A., Jatmiko, S. P., and Rusmiyanto, E. 2015. Dinamika Perubahan Muka Laut Eosen Berdasarkan Data Palinologi Pada Formasi Nanggulan Yogyakarta, *Prosiding Semirata*, 4(1): 442-450.
- Siregar, M. S. 2005. Sedimentasi dan Model Terumbu Formasi Rajamandala di Daerah Padalarang, Jawa Barat, *RISET – Geologi dan Pertambangan*, 15(1): 61-81.
- Slater, S. M., and Wellman, C. H. 2015. A Quantitative Comparison of Dispersed Spore / Pollen and Plant Megafossil Assemblages from A Middle Jurassic Plant Bed from Yorkshire, UK, *Paleobiology*, 41(4): 640-660.
- Subagja, M. A., Setiadi, D. J., Jurnaliah, L., Syafri, I., Alam, S., and Elfitra, D., 2019.
 Pendekatan Kuantitatif Dalam Penentuan Asosiasi Facies Laut Dalam Formasi Halang Pada Sungai Kaligintung, Jawa Tengah, Bulletin of Scientific Contribution: GEOLOGY, 17(3): 193-204.
- Traverse, A. 2007. Paleopalynology 2nd ed, Netherlands: Springer.

- Vernal, A. D., Guiot, J., and Turon, J. L. 1993. Late and postglacial paleoenvironments of the Gulf of St. Lawrence: marine and terrestrial palynological evidence, *Géographie physique et Quaternaire*, 47(2): 167-180.
- Wibowo, E. P., Syafrizal, and Susanto, D. 2016. Jenis Tumbuhan Sumber Nektar Lebah apis dorsata fabr. Dari Desa Bumi Harapan dan Desa Bukit Raya Kecamatan Sepaku Kalimantan Timur, *Bioprospek*, 11(1): 54-64.
- Wibowo, U. P., and Kapid, R. 2014. Biostratigrafi Nannoplankton Daerah Rajamandala, *Jurnal Geologi Sumber Daya Mineral*, 15(4): 185-194.
- Winantris, Syafri, I., and Rahardjo, A. T. 2012. Oncosperma Tigillarium Merupakan Bagian Palino Karakter Delta Plain di Delta Mahakam, Kalimantan, *Bionatura -Jurnal Ilmu-ilmu Hayati dan Fisik*, 14(3): 228-236.
- Winantris, Syafri, I., and Rinawan, R. 2006. Kandungan Mikrofosil Dalam Formasi Pembawa Batubara Dari Daerah Perian Kecamatan Muara Muntai, Kabupaten Kutai, Kertanegara, Kalimantan Timur, Bulletin of Scientific Contribution, 4(1): 8-18.
- Xu, Q., Zhang, S., Gaillard, M., Li, M., Cao, X., Tian, F., and Li, F., 2016. Studies of Modern Pollen Assemblages for Pollen Dispersal - Deposition - Preservation Process Understanding and For Pollenbased Reconstructions of Past Vegetation, Climate, and Human Impact: A Review Based on Case Studies in China, *Quaternary Science Reviews*, 149: 151-166.
- Yulianto, E., Sukapti, W. S., and Setiawan, R.
 2019. Palinostratigrafi, Paleoekologi, dan Paleoklimatologi Plistosen Awal Berdasarkan Studi Palinologi Formasi Pucangan si Daerah Sangiran, Jurnal Geologi dan Sumberdaya Mineral, 20(3): 133-141.